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Reconstruction of geomagnetic activity and near-Earth interplanetary conditions over the past 167 yr – Part 3: Improved representation of solar cycle 11

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Abstract. Svalgaard (2014) has recently pointed out that the calibration of the Helsinki magnetic observatory's H component variometer was probably in error in published data for the years 1866–1874.5 and that this makes the interdiurnal variation index based on daily means, IDV(1d), (Lockwood et al., 2013a), and the interplanetary magnetic field strength derived from it (Lockwood et al., 2013b), too low around the peak of solar cycle 11. We use data from the modern Nurmijarvi station, relatively close to the site of the original Helsinki Observatory, to confirm a 30 % underestimation in this interval and hence our results are fully consistent with the correction derived by Svalgaard. We show that the best method for recalibration uses the Helsinki $Ak(H)$ and aa indices and is accurate to $\pm 10\%$. This makes it preferable to recalibration using either the sunspot number or the diurnal range of geomagnetic activity which we find to be accurate to $\pm 20\%$. In the case of Helsinki data during cycle 11, the two recalibration methods produce very similar corrections which are here confirmed using newly digitised data from the nearby St Petersburg observatory and also using declination data from Helsinki. However, we show that the IDV index is, compared to later years, too similar to sunspot number before 1872, revealing independence of the two data series has been lost; either because the geomagnetic data used to compile IDV has been corrected using sunspot numbers, or vice versa, or both. We present corrected data sequences for both

the IDV(1d) index and the reconstructed IMF (interplanetary magnetic field). We also analyse the relationship between the derived near-Earth IMF and the sunspot number and point out the relevance of the prior history of solar activity, in addition to the contemporaneous value, to estimating any “floor” value of the near-Earth interplanetary field.

Keywords. Geomagnetism and palaeomagnetism (time variations, secular and long term; instruments and technique) – interplanetary physics (interplanetary magnetic fields)

1 Introduction

This paper employs a number of different geomagnetic activity and sunspot indices which are listed, and briefly described, in Appendix A. A review of the reconstruction of conditions in the solar corona and heliosphere from geomagnetic activity was recently presented by Lockwood (2013): a central assumption of all such reconstructions is that a geomagnetic index has, in the past, always responded to varying interplanetary conditions in the same way as it has been observed to do during the space age. Consequently, Lockwood et al. (2013a) compiled the interdiurnal variation geomagnetic index based on daily means, IDV(1d), with the aim of making the construction as homogeneous as possible, such that its response to variations in the near-Earth interplanetary

magnetic field B at all times before the space age was as similar as possible to that in modern times. Lockwood et al. (2013b) subsequently used this index to reconstruct B with the first full analysis of reconstruction errors, carried out using a Monte Carlo technique. This homogeneous construction was achieved by using just three magnetic observatories, in the same latitude band and the same longitude sector, to generate the IDV(1d) composite and by using a model of the secular change in the geomagnetic field to allow for the drift of the observatories in geomagnetic coordinates. The problem with this approach is that unknown errors in the data from a station used would be reproduced in the composite. As a check against this possibility, Lockwood et al. (2013a) compared the IDV(1d) composite against similarly derived values from other nearby stations. This check was necessarily poorer for the data from the 19th century when fewer stations were available for comparisons. Specifically, agreement was found to be very good in comparisons with data from Wilhelmshaven (commencing 1883), Parc St Maur (commencing 1883) and Ekaterinburg (commencing 1887). In addition, after 1880 agreement was shown to be close with the IDV index, derived from a global network of a variable number of observatories (Svalgaard and Cliver, 2010). Before 1863, the IDV(1d) composite data were compared with observations from St Petersburg and the agreement was again very good. Between 1863 and 1883 the only two data sets available for comparisons were from Greenwich and Bartels' interdiurnal daily mean variability index u which, like IDV, was compiled from a variable mix of stations. Agreement between IDV(1d) and u was good after about 1880 but Lockwood et al. (2013a) noted that before then it was partly based on Greenwich data for which both they and Bartels had noted considerable stability problems and could only be readily corrected for temperature variations in a statistical manner. In addition, Lockwood et al. (2013a) were concerned about the u index before 1872 because it was based on diurnal range proxies and so was not a homogeneously constructed index. As a result, despite good correlations with IDV over a period of overlap, it could have different general responses to solar wind and IMF (interplanetary magnetic field) properties.

This concern was shown to be well founded. Because of the stability issues with the Greenwich data, Bartels gave them half weighting and so his interdiurnal u data largely came from the Colaba Observatory, Bombay (now Mumbai), and daily mean interdiurnal variation data for modern stations at such a low geomagnetic latitude station show a $(BV_{\text{SW}})^n$ dependence on solar wind speed V_{SW} with an estimate of n of -0.4 ± 0.9 (Lockwood et al., 2013a), the uncertainty being at the 95 % significance level, derived by applying Fisher's Z test to the correlograms with n (as described by Lockwood, 2002). A result of a negative n is that u would be depressed in intervals of enhanced solar wind speed and this is particularly clear during the declining phases of solar cycles 9 and 10 when u is compared to IDV(1d) which was designed to depend on the IMF strength only (i.e. $n \approx 0$).

This was confirmed to be the case for cycle 9 by the IDV(1d) variation derived from the St Petersburg data, IDV(1d)_{SPE} (the three-letter IAGA observatory code for St Petersburg is SPE). In the light of this finding, Lockwood et al. (2013a) concluded that the difference between u and IDV(1d) at the peak of solar cycle 11 could have arisen for a lower than average solar wind speed at that time and so was not necessarily due to an instrumental error in either data series.

However, Svalgaard (2014) noted that, although the diurnal range in annual means of the Helsinki horizontal (H) component data was similar to that from other stations for most of years for which that station operated (giving usable annual means for 1845–1897), for 8.5 of those years (1866–1874.5) it was lower. Furthermore, using a linear regression of group sunspot number with IDV(1d), values were found to be low at this time, whereas for all other years the agreement was considerably better. He concluded that poor calibration of the “horizontal force” variometer at Helsinki was causing the H component (and hence the IDV(1d) values based on it) to be low in these years. Previously, Nevanlinna (2004) had noted that the range index $Ak(H)_{\text{HLS}}$, scaled from the three-hourly range k index derived from the H data from Helsinki (the IAGA code for Helsinki is HLS) were lower than the linearly regressed aa range index (derived from a Northern Hemisphere and a Southern Hemisphere station) at this time which could also have been caused by the same calibration problem. Nevanlinna (2004) also noted that $Ak(H)_{\text{HLS}}$ at this time was also systematically lower than $Ak(D)_{\text{HLS}}$ (scaled from the station's three-hourly range k index derived from the declination data, D). The H data require a temperature correction and this was implemented by Nevanlinna and Ketola (1993) and, in the absence of any data to check against, a constant sensitivity of the H variometer was applied throughout. These authors also noted that, because the D measurements require no temperature correction and employed a simpler observing geometry, they may be the more reliable of the two in the historical data. From his analysis, Svalgaard (2014) proposed that the Helsinki H , $Ak(H)_{\text{HLS}}$ and IDV(1d) data all need recalibration for the period 1866–1874.5 with an increase of around 30 %. Lockwood et al. (2013a, b) were reluctant to make such a recalibration for a number of reasons: (1) there was considerable variability between $Ak(H)_{\text{HLS}}$, $Ak(D)_{\text{HLS}}$ and aa in all intervals; (2) using sunspot numbers to recalibrate destroys the independence of the geomagnetic and sunspot data sequences; and (3) if the uncertainties in the recalibration were sufficiently large, they could make the IDV(1d) variation appear unwarrantedly similar to the group sunspot data record. This last concern is shown to be well founded in the next section. Subsequent sections study the accuracy and validity of the proposed cycle 11 recalibrations.

2 Analysis of the relationship of IDV and IDV(1d) to sunspot number

Figure 1 presents an analysis of the dependence of the annual means of the geomagnetic indices IDV (right-hand panels) and IDV(1d) (left-hand panels) on the simultaneous annual means of the international sunspot number, R . For IDV(1d), results for before 1872 are given for both the original index as presented in Lockwood et al. (2013a, b; hereafter “Paper 1” and “Paper 2”) (black line) and using the correction derived in the present paper (green line). The top panels (a and b) show the correlations with R^n , as a function of the exponent n . The data have been divided into three intervals: 1844–1872, 1873–1945 and 1946–2012 (inclusive), shown by the black, red and blue lines, respectively. The division on 1 January 1946 was implemented because there has been discussion about a potential discontinuity in R around this date (e.g., Svalgaard, 2012), the division on 1 January 1872 occurs because after this date IDV and IDV(1d) agree closely but less so than before then (see Paper 1). It can be seen that after 1872, both indices agree quite closely on the optimum n value and the level of correlation it yields (the red and blue lines).

The fact that correlations are almost as high for IDV(1d) as for IDV is, at first sight, surprising because IDV is based on a number of stations that varies with date from 1 to over 60, whereas IDV(1d) is based on just one station at any one time. However, IDV uses just one hourly mean value per day, whereas IDV(1d) uses all 24. If we regard data taken by the same station at different Universal Time (UT) as differing only in their UT-dependent calibration, we have a suppression of geophysical and instrument noise by an additional factor in IDV(1d) of $24^{0.5} = 4.89$, which is only achieved using 24 stations for IDV. Hence although noise suppression in IDV(1d) is better for later years (when more than 24 stations are available), it is poorer in early years. In addition, as discussed in Paper 1, the point of IDV(1d) is that it is, unlike IDV, homogeneously constructed to ensure its responses to solar wind parameter variations were as similar in past epochs to those in the space age as it is possible to make them.

The peak correlations r in Fig. 1 for post-1872 data are all in the range 0.821–0.848 and for n in the range 0.5–0.75. The middle panels (c and d) show the variation of the significance S of the difference between the peak correlation and that at general n , as determined by Lockwood (2002). It can be seen that the differences between the $S(n)$ variations are not statistically significant after 1872. However, for neither IDV(1d) or IDV is the same true before 1872. For (uncorrected) IDV(1d), the peak r is 0.806 at $n = 0.29$ (the black line in Fig. 1a). For IDV, however, the peak r is 0.942 at $n = 1.18$ (the black line in Fig. 1b). These discrepancies could denote problems with either the geomagnetic indices or with the sunspot number record, or both. However, it is noticeable that the error is in the opposite sense for the

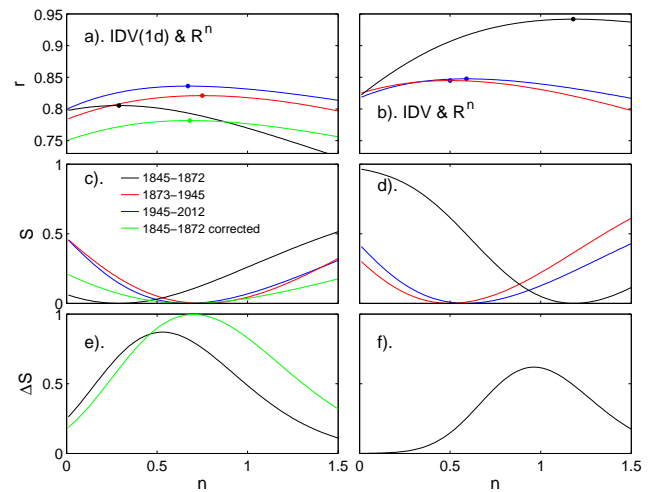


Fig. 1. The dependence of the IDV(1d) (corrected and uncorrected, left panels) and IDV (right panels) on international sunspot number, R , to the power n , as a function of the exponent n . (a) and (b) show the correlation coefficients for the interval 1845–1872 (black line), 1873–1945 (red line) and 1946–2012 (blue line). In the case of IDV(1d) the black line is uncorrected and the green line shows the effect of implementing the correction derived in the present paper. The coloured circle indicates the peak of each correlogram. (c) and (d), the significance S of the differences between the peak correlations and those at general n . (e) and (f), the probability ΔS that the dependence on R before 1872 is the same as after that date (see text for details).

two geomagnetic indices, with peak correlation at somewhat lower n for IDV(1d) but considerably higher n for IDV. The black lines in Fig. 1e and f show the probabilities ΔS that the true n is the same before 1872 as after it: for IDV(1d) this peaks at 87 % whereas for IDV it peaks at just 62 %. Figure 1 therefore shows that neither index is behaving consistently before 1872. In the case of IDV, the correlation with R is far higher than for after 1872 (peaking at 0.942) and is for an optimum value of n that is close to unity. Hence it is clear that the sunspot number and the IDV geomagnetic index are no longer independent data before 1872, making IDV far too similar to R in this interval. This may be because the sunspot number has been used to correct and/or calibrate the geomagnetic data, or vice versa, or both. Hence the concerns about using R to correct geomagnetic indices expressed in Paper 2 are well founded. That having been said, the IDV(1d) index before 1872 is also not behaving in the same way with R as it does in later years, indicating that it too is in error. In the present paper, we derive a correction that does not depend on sunspot number, thereby avoiding the inhomogeneity that Fig. 1 shows that the IDV index suffers from. For completeness, the green lines in Fig. 1a, c and d show the results for IDV(1d) after that correction has been applied. In this case, the n giving a peak correlation of 0.68 is not significantly different to that for the intervals after 1872. The peak correlation coefficient r is lower (0.782), which is to be expected

given the increased uncertainties in both the geomagnetic and sunspot data for the earliest years. The green line in Fig. 1e shows that the probability that the behaviour with R for the corrected IDV(1d) is different before 1872 to that after 1872 is just 0.003.

3 Retrospective recalibration techniques

3.1 Using the diurnal range of geomagnetic data from other observatories

Svalgaard (2014) makes use of the measured diurnal variation in annual means of the H component to recalibrate the H variometer. He points out that the average diurnal range (the difference between the maximum and minimum values of the average diurnal variation, over an extended period such as a year) is very well correlated with solar activity indices such as group sunspot number. That being the case, the average diurnal range should be the same at all stations (provided they are removed from the influence of auroral currents) and so comparing measured diurnal ranges from different stations allows one to intercalibrate those stations and comparing to group sunspot number provides a second test. However, the term “very well correlated” is subjective and does not quantify the uncertainty inherent in adopting this approach. Correlation coefficients r are typically 0.9 meaning that $(1-r^2) \approx 0.2$ of the variation at one station is not matched at another and hence calibration using these correlations will typically be accurate to about 20 %.

The red and green lines in Fig. 2 show the annual mean diurnal range of the H component observed by modern-day stations Nurmijarvi and Eskdalemuir (ΔH_{NUR} and ΔH_{ESK}). Nurmijarvi is the closest magnetometer to the old Helsinki Observatory and Eskdalemuir and Helsinki are the primary observatories contributing to the IDV(1d) index (the coordinates of all three stations are given in Table 1 of Paper 1). Also shown (in black) is the linearly regressed group sunspot number, R_G , as observed by the network coordinated by the Royal Greenwich Observatory (RGO) until 1976 and, thereafter, from the US Air Force (USAF) and US National Oceanic and Atmospheric Administration (NOAA) Solar Optical Observing Network (SOON). These data have been homogenised into a single data set (see Hathaway, 2010). The correlations are indeed good, between ΔH_{NUR} and ΔH_{ESK} it is 0.972 (significant at >99.999 % level by comparison against the AR1 red-noise model), between ΔH_{NUR} and R_G it is 0.897 (significant at the 75 % level) and between ΔH_{ESK} and R_G it is 0.935 (significant at the 78 % level). However it can be seen that there are differences between the two diurnal range variations, potentially due to site differences but also due to calibration uncertainties. In particular, agreement between the diurnal ranges and sunspot numbers during solar cycle 20 (peaking around 1970) is notably poor, especially if

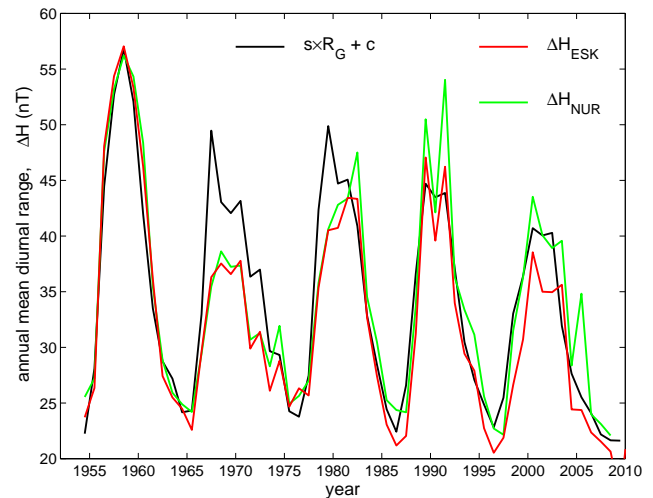


Fig. 2. The annual, mean diurnal range for modern day stations Nurmijarvi (in green) and Eskdalemuir (in red), ΔH_{NUR} and ΔH_{ESK} for the period 1953–2008. Also shown (in black) are the best fit linear regression to ΔH_{NUR} of the annual means of the group sunspot number R_G taken from the Royal Greenwich Observatory observations up to 1976 and the USAF/NOAA SOON network thereafter.

group sunspot numbers, R_G , are used rather than the international sunspot number, R .

Figure 3 shows the best linear regression fits of the diurnal range variations from a large number of stations (in grey). These were evaluated for a total of 81 stations. However it was found auroral stations correlated badly and so all stations poleward of the corrected geomagnetic latitude $\Lambda = 60^\circ$ were excluded. This left 67 stations and of these a further 5 were removed because the correlation with ΔH_{NUR} was below 0.65, indicating possible calibration problems. The green and mauve lines show the variations of the 5th and 95th percentiles of the distribution for each year and it can be seen that there is considerable spread. Some of this arises from intercalibration problems, but there may also be site changes, noise interference and genuine differences in the behaviour at different sites: in particular, sites near the upper limit of $\Lambda = 60^\circ$ will be subject to greater auroral effects and this includes Nurmijarvi itself (corrected geomagnetic latitude $\Lambda = 56.91^\circ$).

The top panel in Fig. 4 presents the total distribution of the fit residuals inherent in the fits shown in Fig. 3, for all years and all 62 stations, as a percentage of the simultaneous ΔH_{NUR} value. The vertical red lines mark the 5th and 95th percentiles which are at percentage deviations of -19% and $+19\%$. Thus even for modern stations, using data from one other station gives a calibration that is accurate for any one annual mean to $\pm 19\%$ (at the 2σ level). This value is consistent with the expectations from the magnitude of the correlation coefficients. For the 19th-century data, there are very few stations available and if their data have been, at any point, intercalibrated using the diurnal range, these stations

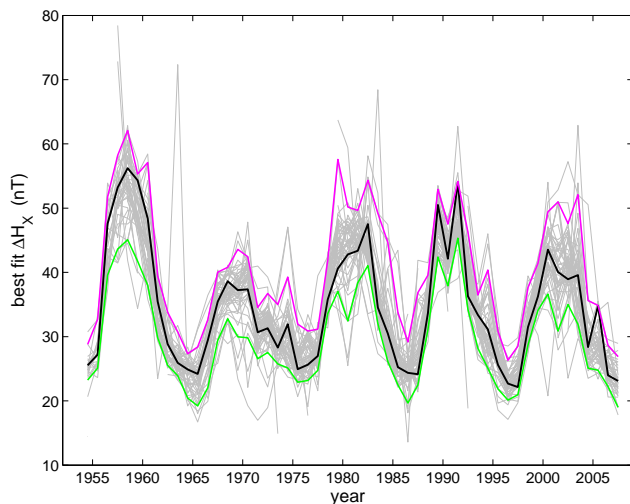


Fig. 3. Annual, mean diurnal range variations from 62 mid- and low-latitude stations (in grey), ΔH_X (where X is the observatory code), linearly regressed against the variation for Nurmijarvi (in black), ΔH_{NUR} . The green and mauve lines are the 5th and 95th percentiles of the distribution for each year. The 62 stations selected from the total of 81 analysed are those that gave a correlation with ΔH_{NUR} exceeding 0.65 and were at a corrected magnetic latitude Λ below 60° .

will not be providing independent calibration. It is important to note that, because it is at a higher latitude, there is an auroral component to the response of Helsinki, indeed that response is important in the development of IDV(1d) as it gives a dependence on IMF B rather than the $BV^{-0.4}$ of lower latitude stations (see Paper 1). Thus at least some of the $\pm 19\%$ uncertainty arises from the latitude of Helsinki.

Importantly, Svalgaard's (2014) estimate of a 30 % error during the period 1866–1874.5 is outside the $\pm 19\%$ uncertainty inherent in the recalibration procedure he adopted – confirming the Helsinki data are indeed too low in this interval and potentially correcting the error in IDV(1d) also identified in Fig. 1.

3.2 Using group sunspot numbers

Svalgaard (2014) also tests diurnal range estimates against group sunspot numbers. However, Fig. 2 shows that although the correlation is high, the inherent errors in this recalibration can also be quite high. The middle panel in Fig. 4 shows the distribution of percentage residuals obtained by fitting the Nurmijarvi and then the Eskdalemuir data with group sunspot number. The largest error seen is again -19% . Note this is a much smaller sample than in the top panel, using just two data series rather than 63, however, the overall error appears to be similar to that obtained in the previous section.

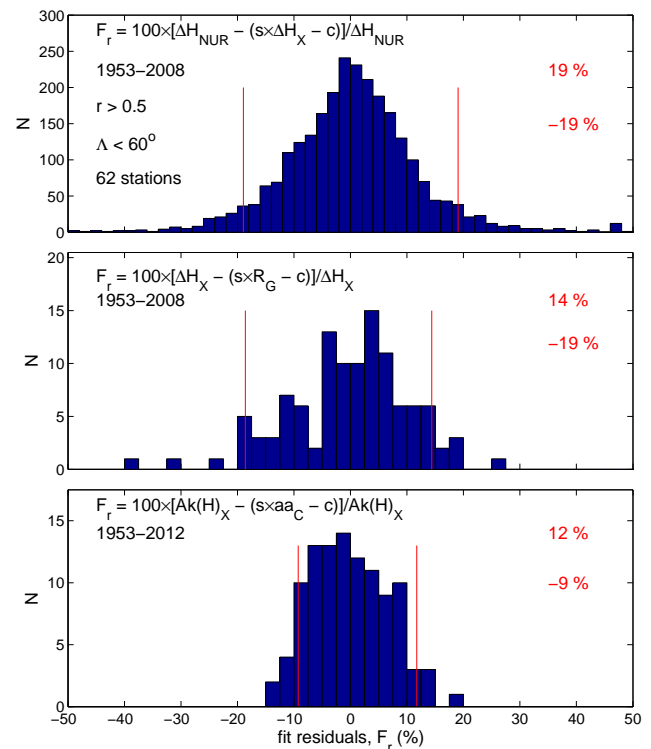


Fig. 4. Distributions of percentage, linear-regression fit residuals. (Top panel) For the linear regression fits of ΔH_X (where X is the observatory code), to ΔH_{NUR} for the 62 variations shown in Fig. 3. (Middle panel) For the linear regression fits of group sunspot number to the ΔH_{NUR} and ΔH_{ESK} variations. (Bottom panel) For the linear regression fits of the aa range index to $Ak(H)_{NUR}$ and $Ak(H)_{ESK}$. In each panel the vertical, red lines give the 5th and 95th percentiles of the distribution, the values of which are given in red.

3.3 Using geomagnetic three-hourly-range k values

Figure 5 compares the $Ak(H)$ series obtained from the range k values from the horizontal field component H measured at Eskdalemuir and Nurmijarvi with the aa_C index, also derived from k values. The aa_C index is based on the standard aa index of Mayaud (1971, 1972, 1980) with some corrections devised by Lockwood et al. (2006) by comparison with range k indices and/or the Ap index based on the k indices from a network of midlatitude stations. The biggest difference between aa and aa_C is a 2.5 nT shift around 1957 associated with the move of the Northern Hemisphere's aa station from Abinger to Hartland. The correction to aa to yield aa_C is discussed further in Paper 4 (Lockwood et al., 2014). The correlation between $Ak(H)_{NUR}$ and $Ak(H)_{ESK}$ is 0.948 (significant at the 85.3 % level), between $Ak(H)_{ESK}$ and aa_C it is 0.980 (95 %) and between $Ak(H)_{NUR}$ and aa_C it is 0.975 (91 %). From these levels of correlation we would expect the use of $Ak(H)$ and aa_C to give calibrations that are accurate to the order of 5–10 %.

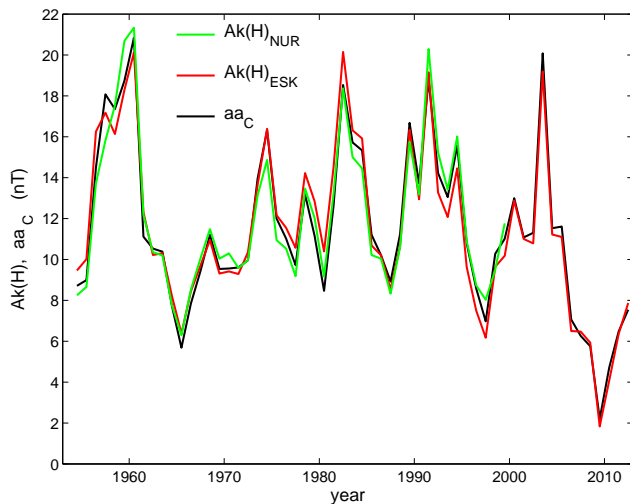


Fig. 5. The variations of the $Ak(H)$ indices derived from the horizontal field observed at Nurmijarvi, $Ak(H)_{\text{NUR}}$ (green, available until 2001), and Eskdalemuir $Ak(H)_{\text{ESK}}$ (red) and the linearly regressed corrected aa_C (black).

The bottom panel in Fig. 4 shows the distribution of percentage fit residuals of the linear regression of aa_C to $Ak(H)_{\text{NUR}}$ and $Ak(H)_{\text{ESK}}$. As in the other two panels, the vertical red lines mark the 5th and 95th percentiles which are at -9% and $+12\%$. This spread is close to half that found using the diurnal range and the group sunspot number. Thus we find comparing with aa to be preferable to comparing diurnal ranges or diurnal range with group sunspot number.

Looking at the 19th-century data from Helsinki, the correlation between $Ak(H)_{\text{HLS}}$ and aa_C for the interval of available data (1868–1897, inclusive) is 0.85 (significance $> 99.99\%$). However, this contains the interval for which Svalgaard (2014) finds the calibrations of the horizontal component instrument to be poorly calibrated (1866–1874.5), and removing these data causes the correlation to rise to 0.92 (however the significance falls to 99 % due to the lower number of data points). The best-fit linear regression for this second fit is

$$[Ak(H)_{\text{HLS}}]_{\text{fit}} = 0.843 \times aa_C - 0.039. \quad (1)$$

Note that for the modern Nurmijarvi data (1953–2011) the corresponding slope is 1.026 with an intercept of zero to within seven decimal figures. The ratio of the regression slopes with aa_C for the modern Nurmijarvi data and the historic Helsinki data (0.843/1.026) implies that the sensitivity of the 19th century Helsinki instrument and this is consistent with the intercalibrations used to generate IDV(1d), as described in Paper 1. The ratio $[Ak(H)_{\text{HLS}}]_{\text{fit}}/Ak(H)_{\text{HLS}}$ from Eq. (1) gives us a correction factor which we can apply to the data from 1868 (when aa_C is available) and Fig. 4c indicates that this recalibration can be expected to be accurate to within about $\pm 10\%$.

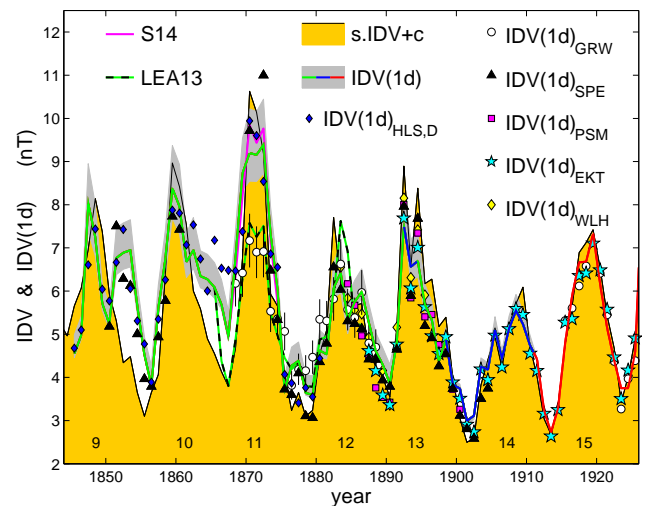


Fig. 6. Revised version of Fig. 14 of Lockwood et al. (2013a) (Paper 1). Annual means of the IDV(1d) composite are shown with data originating from ESK in red, NGK in blue and HLS in green, surrounded by a grey area showing the band of uncertainty that arises from the two regression fits. The HLS data shown by the green line have been corrected in cycle 11 as discussed in the text. The black line bounding the filled orange area is the IDV index, scaled using the regression for the period 1880–2013, as shown in Fig. 12 of Lockwood et al. (2013a). IDV(1d) derived from various early data sets are also shown: from St Petersburg, IDV(1d)_{SPE} (black triangles); from Ekaterinburg, IDV(1d)_{EKA} (cyan stars); from Parc St Maur, IDV(1d)_{PSM} (mauve squares); from Wilhelmshaven, IDV(1d)_{WLH} (yellow diamonds); and from Greenwich, IDV(1d)_{GRW} (white circles). The blue diamonds are derived using the D component observed at Helsinki, IDV(1d)_{HLS,D}. The Greenwich values are compiled using the daily means of H that are uncorrected for temperature variations, with annual IDV(1d)_{GRW} means subsequently corrected (giving the shown uncertainty bands caused by the temperature effects). Solar cycle numbers are shown across the base of the figure. The green-and-black dashed line in Fig. 6 is the variation of IDV(1d) derived Paper 1 from the uncorrected Helsinki magnetometer data. The mauve line shows the data series with the 30 % upward recalibration for the period 1866–1874.5 derived by Svalgaard (2014).

4 The corrected IDV(1d) data series

The green-and-black dashed line in Fig. 6 is the variation of IDV(1d) derived in Paper 1 from the uncorrected Helsinki magnetometer data. The mauve line shows the data series with the 30 % upward recalibration for 1866–1874.5 derived by Svalgaard (2014). The green line uses the correction described in Sect. 2.3, derived from the observed aa_C index for the period 1868–1899, which yields $[Ak(H)_{\text{HLS}}]_{\text{fit}}$ from Eq. (1) and hence the correction factor to H values $[Ak(H)_{\text{HLS}}]_{\text{fit}}/Ak(H)_{\text{HLS}}$ from Eq. (1) (which Fig. 4c shows to be accurate to within $\pm 10\%$). This correction cannot be applied before the start of the aa_C data in 1868 and for the period 1866–1867, the correction of Svalgaard (2014), based

on diurnal range, is applied and in Fig. 4a is shown to be accurate to within $\pm 20\%$. The green and mauve lines are particularly close for the period 1868–1870 which means that this procedure does not introduce a discontinuity into the correction.

5 Comparison with St Petersburg and declination data

We have recently completed digitising hourly means of H component data in the yearbooks from the St Petersburg observatory (Kupffer, 1853–1865; Wild, 1872–1895; Rykatchew, 1896–1898, 1899–1908). The black triangles in Fig. 6 show the annual means derived from daily means of these data, $IDV(1d)_{SPE}$, presented here for the first time. Absolute H -component data are available for 1870 and 1873–1906, inclusive. Note that in Fig. 6, no scaling of the $IDV(1d)_{SPE}$ data for 1870 and after has been carried out. After 1883, agreement is very good with other data from Wilhelmshaven, Parc St Maur and Ekaterinburg ($IDV(1d)_{WLH}$, $IDV(1d)_{PSM}$, and $IDV(1d)_{EKA}$ – shown by the yellow diamonds, mauve squares and cyan stars, respectively). These data have also not been scaled. The good agreement of all these stations with the Helsinki data ($IDV(1d)_{HLS}$ – green line) is good confirmation that the fit of the calibrations derived in Paper 1 to the Eskdalemuir data (using the Niemegk data) is broadly correct. Note that the earlier St Petersburg data (for up to 1863) has required an upward rescaling of 11 % to fit the amplitude of the variation seen in the Helsinki data. This factor is derived from comparing the $Ak(H)$ data for Helsinki and St Petersburg (Nevanlinna and Häkkinen, 2010), but we have, at least up to the time of writing, no absolute independent calibration of either the Helsinki or St Petersburg H variometers at this time.

As a further test we have computed $IDV(1d)$ using the declination data from Helsinki. The results for $IDV(1d)_{HLS,D}$ are shown by the blue diamonds in Fig. 6. These data have been scaled by linear regression to the standard $IDV(1d)$ data from Helsinki (which, using the same notation, would be $IDV(1d)_{HLS,H}$) to allow for the differences between H and D . Figure 6 shows that the corrected H and D values are give very similar variations at all times. The worst deviations being for the period 1865–1867, for which the correction is most uncertain.

During solar cycles 12 and 13, the various stations are all giving broadly similar results. All the stations imply IDV is slightly too high during the declining phase of cycle 13, but all the data agree well on the first peak of cycle 13 (1892). The station data suggest that the second peak (1894) is slightly too large in IDV but slightly too small in $IDV(1d)$. Agreement of all stations and with both IDV and $IDV(1d)$ is better in cycle 12. The correction of the Helsinki data using a_{AC} has slightly reduced the peak of this cycle in $IDV(1d)$ so it now agrees better with the data from the other stations. The white circles show the Greenwich data, and the relatively

good agreement found in this cycle encouraged Lockwood et al. (2013a) to think it may be useful, for all its limitations, in cycle 11. Figure 6 shows us that this is not the case and the Greenwich values in cycle 11 are too low. It is not clear if this is because of an over-correction for the stability problems, because of the statistical temperature correction used or because of changes in the observing procedure and/or equipment at Greenwich.

The poor reliability of the Greenwich data in cycle 11, calls into question its use as a check on the variation between cycles 11 and 12. This being the case, the St Petersburg data are the best test and support the slightly deeper minimum in IDV . At the peak of cycle 11, the available $IDV(1d)_{SPE}$ data (1870 and 1873) are closer to the Svalgaard correction than the one implemented here, but given they are intrinsically accurate to about $\pm 20\%$ and $\pm 10\%$, respectively, these differences are not significant and either variation is equally valid.

For cycles 9 and 10, the St Petersburg data provide a test of the Helsinki data. Svalgaard notes the Helsinki H data are consistent with the group sunspot data but, as shown in Fig. 4b, this test is only accurate to 20 %. The St Petersburg data match well after rescaling with a linear regression factor of 1.1 which is well within the uncertainty of the sunspot test.

The St Petersburg data show the major declining phase peak in cycle 9 is real. This peak is absent completely in IDV which, at this time, is based on a best-fit diurnal proxy and not on interdiurnal variation. This strong suppression appears to be due to an inverse dependence of the proxy on solar wind flow speed. This being the case, the diurnal-proxy u data used to extend the IDV into cycles 8–11 (Svalgaard and Cliver, 2010) must be used with great caution and recognising that it does not have the dependence on B alone which is the important characteristic of interdiurnal variation indices.

6 Revised reconstruction of the near-Earth IMF

We have reapplied the analysis of Lockwood et al. (2013b) (Paper 2) to the corrected $IDV(1d)$ to reconstruct the IMF B from 1845 onwards. The result is shown in Fig. 7. The dark grey band shows the uncertainty associated with the regression and the joins between data from different stations, and so is the corrected version of the error band shown in Fig. 7 of Paper 2. The light grey area in Fig. 7 of the present paper shows this uncertainty combined with that in the raw data from which $IDV(1d)$ is compiled, obtained by comparing $Ak(H)$ and $Ak(D)$ with data from nearby stations (for the earliest data this is from St Petersburg only). The major difference from the variation presented in Paper 2 is that the peak of cycle 11 is considerably higher such that reconstructions based on $IDV(1d)$ and IDV are no longer different at this time to within the uncertainties. In addition, the minimum before cycle 11 is much less deep, which is consistent with the cycle to cycle growth of open solar flux, reflected in the near-Earth IMF, which gave the large peak in cycle

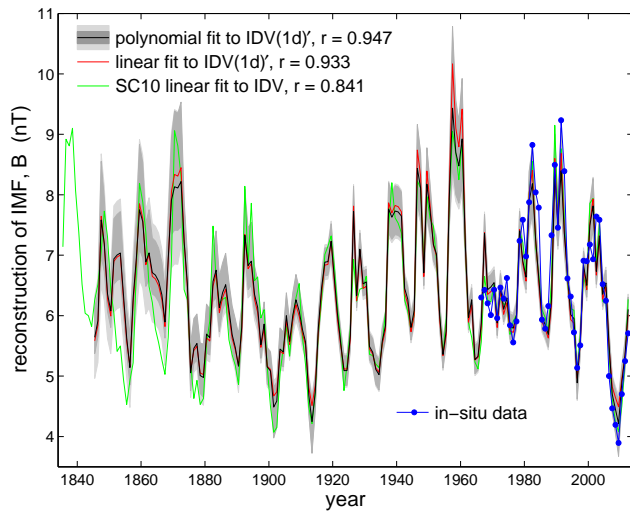


Fig. 7. Corrected version of Fig. 7 of Lockwood et al. (2013b) which gives reconstructions of the near-Earth IMF, B , from geomagnetic data. The black line uses the corrected geomagnetic activity composite, IDV(1d) and the polynomial fit to B derived in Lockwood et al. (2013b). The dark grey area surrounding this black line is the uncertainty band associated with using this polynomial fit derived using a Monte Carlo technique (see Lockwood et al., 2013b) and the light grey area convolves this uncertainty with the measurement uncertainty. The red line shows the best reconstruction using their linear fit. The green line shows the SC10 reconstruction. Blue dots show the annual means of the observed IMF.

11. However, we note these values are corrected using the diurnal range method and that correction is only accurate to $\pm 20\%$. Table 1 gives the full, corrected data sequences for IDV(1d) and IMF B with their associated uncertainty bands.

7 Relationship of near-Earth IMF to group sunspot number

Svalgaard and Cliver (2005) note a correlation between the IMF and the square root of the simultaneous group sunspot number (in annual mean data). In this section, we investigate this correlation in the reconstructed IMF shown in Fig. 7. The top panel in Fig. 8 shows correlograms of group sunspot number (to power n), R_G^n , against the near-Earth IMF B for annual mean data. The square of the correlation coefficient r^2 is shown as a function of n from 0 to 3 for (in red) IMF observations (1964–2012, inclusive), (in black) the reconstructed IMF (1845–2012) from the IDV(1d) index, and (in blue) the modelled IMF (1612–2012). The modelled IMF is based on the modelling of open solar flux (OSF) by Owens and Lockwood (2012) from the continuity equation, using the group sunspot number to quantify the OSF emergence rate (Solanki et al., 2000), and a loss rate that depends on the current sheet tilt (Owens et al., 2011). This is converted into B using the empirical relationship by Lockwood et al. (2009)

Table 1. The full data series with corrected data for the period 1866–1875. Annual IDV(1d) composite and derived IMF B values are given with 2σ uncertainties.

Year	IDV(1d) (nT)			IMF B (nT)		
	best	min.	max.	best	min.	max.
1846.5	4.99	4.87	5.40	6.01	5.82	6.37
1847.5	8.04	7.38	8.94	7.78	7.33	8.41
1848.5	7.26	6.79	7.99	7.36	7.02	7.89
1849.5	5.73	5.55	6.16	6.47	6.27	6.82
1850.5	5.25	5.12	5.65	6.18	5.99	6.52
1851.5	6.78	6.37	7.45	7.10	6.78	7.57
1852.5	6.90	6.49	7.57	7.16	6.85	7.64
1853.5	6.96	6.55	7.64	7.20	6.88	7.68
1854.5	5.66	5.45	6.15	6.43	6.21	6.81
1855.5	4.55	4.46	4.93	5.71	5.53	6.08
1856.5	3.83	3.76	4.26	5.21	5.00	5.65
1857.5	5.62	5.39	6.13	6.41	6.17	6.80
1858.5	6.76	6.37	7.42	7.09	6.77	7.56
1859.5	8.39	7.68	9.36	7.97	7.48	8.64
1860.5	7.98	7.31	8.92	7.75	7.29	8.40
1861.5	6.69	6.31	7.34	7.04	6.74	7.51
1862.5	6.97	6.51	7.69	7.20	6.86	7.71
1863.5	6.36	6.07	6.93	6.85	6.60	7.27
1864.5	6.29	5.99	6.84	6.81	6.55	7.22
1865.5	6.10	5.83	6.65	6.70	6.45	7.11
1866.5	5.69	5.57	6.27	6.34	6.08	6.88
1867.5	4.94	4.84	5.54	5.90	5.63	6.46
1868.5	6.77	6.59	7.43	6.94	6.63	7.58
1869.5	8.73	8.34	9.54	7.92	7.34	8.94
1870.5	9.20	8.55	10.22	8.13	7.42	9.39
1871.5	9.15	8.54	10.16	8.11	7.41	9.35
1872.5	9.39	8.73	10.45	8.22	7.47	9.53
1873.5	7.32	7.11	7.95	7.23	6.86	7.91
1874.5	6.02	5.90	6.51	6.53	6.26	7.03
1875.5	3.70	3.65	4.05	5.06	4.76	5.54
1876.5	4.27	4.22	4.70	5.46	5.21	5.95
1877.5	4.39	4.31	4.82	5.54	5.27	6.03
1878.5	3.64	3.59	4.06	5.01	4.71	5.55
1879.5	3.60	3.51	4.04	4.98	4.65	5.53
1880.5	4.62	4.47	5.09	5.69	5.38	6.19
1881.5	4.54	4.43	4.95	5.64	5.35	6.11
1882.5	6.08	5.84	6.63	6.56	6.23	7.10
1883.5	6.42	5.97	7.13	6.75	6.30	7.40
1884.5	5.33	5.01	5.88	6.13	5.74	6.66
1885.5	5.75	5.61	6.25	6.38	6.10	6.87
1886.5	5.99	5.83	6.52	6.51	6.23	7.03
1887.5	5.28	5.20	5.71	6.10	5.85	6.56
1888.5	4.61	4.54	5.04	5.69	5.43	6.16
1889.5	4.29	4.25	4.69	5.48	5.23	5.95
1890.5	3.87	3.78	4.29	5.23	5.02	5.67
1891.5	4.79	4.79	5.03	5.88	5.77	6.14
1892.5	7.51	7.5	8.22	7.50	7.39	8.01
1893.5	6.57	6.2	7.19	6.98	6.67	7.43
1894.5	6.71	6.33	7.36	7.06	6.75	7.52
1895.5	5.71	5.51	6.16	6.46	6.25	6.82
1896.5	5.35	5.2	5.74	6.24	6.05	6.57
1897.5	4.35	4.34	4.52	5.58	5.44	5.82
1898.5	4.88	4.87	5.14	5.94	5.82	6.21
1899.5	3.83	3.78	3.95	5.20	5.02	5.43

Table 1. Continued.

Year	IDV(1d) (nT)			IMF <i>B</i> (nT)		
	best	min.	max.	best	min.	max.
1900.5	3.74	3.69	3.84	5.14	4.95	5.36
1901.5	2.99	2.89	3.01	4.55	4.28	4.74
1902.5	3.11	3.02	3.13	4.65	4.40	4.84
1903.5	4.25	4.19	4.45	5.51	5.33	5.77
1904.5	4.18	4.15	4.33	5.45	5.30	5.69
1905.5	5.10	5.09	5.4	6.08	5.97	6.37
1906.5	4.43	4.42	4.63	5.64	5.50	5.88
1907.5	5.12	5.12	5.41	6.09	5.99	6.37
1908.5	5.54	5.54	5.91	6.36	6.27	6.67
1909.5	5.23	5.23	5.54	6.16	6.07	6.45
1910.5	4.74	4.74	4.98	5.85	5.73	6.11
1911.5	4.43	4.43	4.44	5.63	5.51	5.77
1912.5	3.24	3.24	3.24	4.75	4.58	4.92
1913.5	2.73	2.73	2.73	4.33	4.14	4.52
1914.5	3.29	3.29	3.29	4.79	4.62	4.96
1915.5	4.79	4.79	4.79	5.87	5.76	5.99
1916.5	5.89	5.89	5.89	6.58	6.49	6.66
1917.5	6.67	6.67	6.67	7.03	6.94	7.12
1918.5	6.68	6.68	6.68	7.04	6.95	7.13
1919.5	7.33	7.33	7.33	7.40	7.30	7.50
1920.5	6.17	6.17	6.17	6.74	6.66	6.82
1921.5	5.14	5.14	5.14	6.11	6.01	6.21
1922.5	4.39	4.39	4.39	5.61	5.48	5.73
1923.5	3.75	3.75	3.75	5.15	5.00	5.30
1924.5	3.75	3.75	3.75	5.15	5.00	5.30
1925.5	4.59	4.59	4.59	5.74	5.62	5.86
1926.5	8.32	8.32	8.32	7.93	7.79	8.07
1927.5	5.67	5.67	5.67	6.44	6.35	6.53
1928.5	7.07	7.07	7.07	7.26	7.17	7.35
1929.5	6.01	6.01	6.01	6.65	6.56	6.73
1930.5	6.07	6.07	6.07	6.68	6.60	6.77
1931.5	4.32	4.32	4.32	5.55	5.42	5.68
1932.5	4.17	4.17	4.17	5.45	5.32	5.59
1933.5	3.78	3.78	3.78	5.17	5.02	5.31
1934.5	3.66	3.66	3.66	5.08	4.93	5.23
1935.5	4.45	4.45	4.45	5.65	5.52	5.77
1936.5	4.85	4.85	4.85	5.92	5.81	6.02
1937.5	8.40	8.40	8.40	7.97	7.83	8.12
1938.5	8.12	8.12	8.12	7.83	7.70	7.96
1939.5	8.33	8.33	8.33	7.94	7.80	8.08
1940.5	8.30	8.30	8.30	7.92	7.78	8.06
1941.5	8.16	8.16	8.16	7.85	7.71	7.98
1942.5	5.87	5.87	5.87	6.56	6.48	6.65
1943.5	5.57	5.57	5.57	6.38	6.29	6.47
1944.5	4.80	4.80	4.80	5.88	5.77	5.99
1945.5	5.26	5.26	5.26	6.18	6.08	6.28
1946.5	9.88	9.88	9.88	8.71	8.50	8.92
1947.5	9.17	9.17	9.17	8.36	8.18	8.54
1948.5	6.34	6.34	6.34	6.84	6.76	6.93
1949.5	9.29	9.29	9.29	8.42	8.24	8.60
1950.5	8.32	8.32	8.32	7.93	7.79	8.07
1951.5	7.22	7.22	7.22	7.34	7.25	7.44
1952.5	6.92	6.92	6.92	7.17	7.08	7.26
1953.5	5.10	5.10	5.10	6.08	5.98	6.18
1954.5	4.13	4.13	4.13	5.42	5.28	5.55
1955.5	4.71	4.71	4.71	5.82	5.71	5.94
1956.5	8.83	8.83	8.83	8.19	8.03	8.35

Table 1. Continued.

Year	IDV(1d) (nT)			IMF <i>B</i> (nT)		
	best	min.	max.	best	min.	max.
1957.5	12.29	12.29	12.29	9.80	9.47	10.14
1958.5	10.49	10.49	10.49	8.99	8.75	9.24
1959.5	9.97	9.97	9.97	8.75	8.54	8.97
1960.5	11.02	11.02	11.02	9.24	8.97	9.51
1961.5	7.18	7.18	7.18	7.32	7.23	7.42
1962.5	5.06	5.06	5.06	6.06	5.95	6.16
1963.5	5.55	5.55	5.55	6.36	6.27	6.45
1964.5	4.00	4.00	4.00	5.33	5.19	5.47
1965.5	4.08	4.08	4.08	5.38	5.25	5.52
1966.5	4.91	4.91	4.91	5.96	5.85	6.06
1967.5	7.57	7.57	7.57	7.54	7.43	7.65
1968.5	5.84	5.84	5.84	6.55	6.46	6.63
1969.5	5.92	5.92	5.92	6.59	6.51	6.68
1970.5	6.05	6.05	6.05	6.67	6.59	6.75
1971.5	4.91	4.91	4.91	5.96	5.85	6.07
1972.5	5.94	5.94	5.94	6.61	6.52	6.69
1973.5	5.28	5.28	5.28	6.20	6.10	6.29
1974.5	5.60	5.60	5.60	6.40	6.31	6.49
1975.5	4.82	4.82	4.82	5.89	5.78	6.00
1976.5	4.78	4.78	4.78	5.87	5.76	5.98
1977.5	5.13	5.13	5.13	6.10	6.00	6.20
1978.5	7.66	7.66	7.66	7.58	7.47	7.69
1979.5	6.69	6.69	6.69	7.05	6.96	7.13
1980.5	6.21	6.21	6.21	6.77	6.68	6.85
1981.5	8.68	8.68	8.68	8.12	7.96	8.27
1982.5	9.31	9.31	9.31	8.43	8.25	8.61
1983.5	7.39	7.39	7.39	7.44	7.33	7.54
1984.5	6.20	6.20	6.20	6.76	6.67	6.84
1985.5	5.07	5.07	5.07	6.06	5.96	6.16
1986.5	4.87	4.87	4.87	5.93	5.82	6.04
1987.5	4.62	4.62	4.62	5.76	5.64	5.88
1988.5	5.57	5.57	5.57	6.38	6.29	6.47
1989.5	9.65	9.65	9.65	8.59	8.40	8.79
1990.5	7.83	7.83	7.83	7.67	7.55	7.79
1991.5	9.79	9.79	9.79	8.66	8.46	8.87
1992.5	7.57	7.57	7.57	7.53	7.42	7.64
1993.5	6.35	6.35	6.35	6.85	6.76	6.93
1994.5	5.13	5.13	5.13	6.10	6.00	6.20
1995.5	5.03	5.03	5.03	6.04	5.93	6.14
1996.5	3.48	3.48	3.48	4.94	4.78	5.10
1997.5	4.46	4.46	4.46	5.65	5.53	5.78
1998.5	6.58	6.58	6.58	6.98	6.90	7.07
1999.5	5.97	5.97	5.97	6.62	6.54	6.71
2000.5	7.84	7.84	7.84	7.68	7.56	7.80
2001.5	8.51	8.51	8.51	8.03	7.88	8.18
2002.5	6.38	6.38	6.38	6.87	6.78	6.95
2003.5	7.49	7.49	7.49	7.49	7.38	7.60
2004.5	5.71	5.71	5.71	6.46	6.37	6.55
2005.5	6.04	6.04	6.04	6.66	6.58	6.75
2006.5	4.38	4.38	4.38	5.60	5.47	5.72
2007.5	3.25	3.25	3.25	4.76	4.59	4.93
2008.5	2.95	2.95	2.95	4.52	4.34	4.70
2009.5	2.70	2.70	2.70	4.30	4.11	4.49
2010.5	3.48	3.48	3.48	4.94	4.78	5.10
2011.5	4.12	4.12	4.12	5.41	5.28	5.55
2012.5	5.28	5.28	5.28	6.19	6.10	6.29

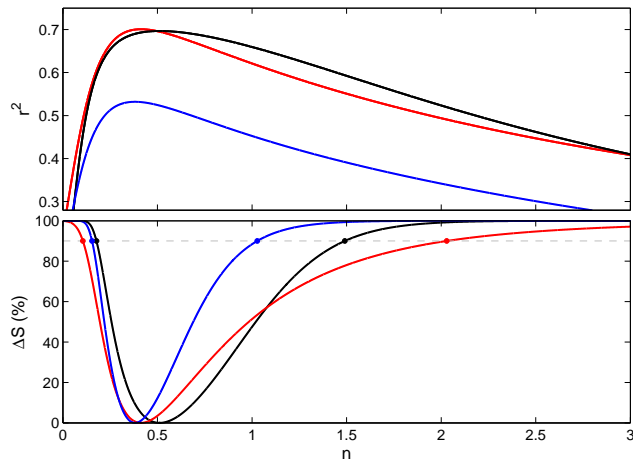


Fig. 8. (Top) Correlograms of group sunspot number, (to power n) R_G^n , against the IMF B for annual mean data. The square of the correlation coefficient r^2 is shown as a function of n from 0 to 3 for: (in red) IMF observations (1964–2012, inclusive), (in black) the reconstructed IMF (1845–2012) from the IDV(1d) index, and (in blue) using modelled IMF (1611–2012). (Bottom) The significance ΔS of the difference between the peak correlation and the correlation at a general n . The horizontal dashed grey line is the 90 % significance level.

and the resulting time series is shown in Fig. 1 of Lockwood and Owens (2014). The plot shows that for the modelled and observed IMF data series, the correlations peak at $n = 0.4$, and for the geomagnetically reconstructed IMF it peaks at $n = 0.5$. The peak correlation coefficients r and their significance S (computed against the AR-1 auto-regressive “red-noise” model) are $r = 0.837$ ($S = 96.3\%$) for the IMF data, $r = 0.835$ ($S = 99.98\%$) for the geomagnetic reconstruction and $r = 0.729$ ($S > 99.99\%$) for the modelled data. Hence for the observations and geomagnetic reconstruction of the IMF, R_G^n explains about 70 % of the IMF variation, but R_G^n only explains 53 % of the modelled variation. The bottom panel in Fig. 8 shows the significance ΔS of the difference between the peak correlation and the correlation at a general n using Fisher’s Z test as described by Lockwood (2002). The dashed grey line is the 90 % significance level and at this confidence level the uncertainty in the peak n is lower for the longer data series. The length of the data series means that the upper limits to n are 2, 1.5 and 1 for the IMF, geomagnetic and model data series; for all three cases the lower limit is near $n = 0.2$. Given this large uncertainty in the peak n estimates (at the 90 % confidence level), the difference in the behaviour of the reconstructed IMF and the other two is not significant. Figure 9 explains why the correlation for the modelled sequence is lower. This is a scatter plot of B as a function of $R_G^{0.4}$ for the three data series: the in situ observations are shown by red triangles, the reconstructed IMF by black dots (with uncertainties in B as shown in Fig. 7), the modelled IMF since 1845 by cyan squares and the modelled IMF before 1845 by yellow diamonds. The observations, the reconstruction and

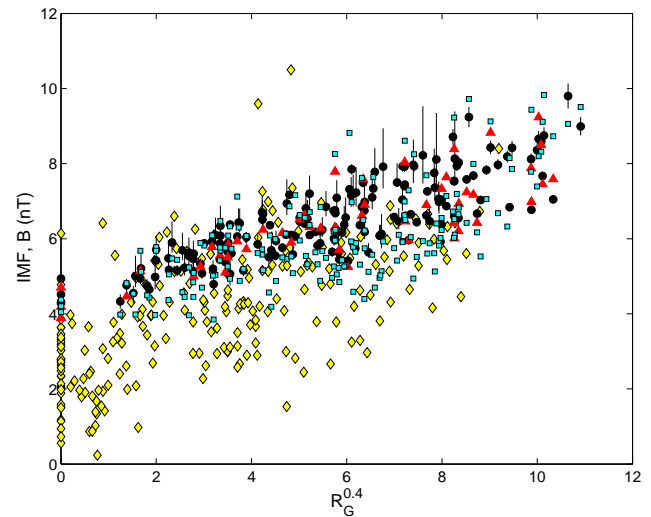


Fig. 9. Scatter plot of near-Earth IMF B as a function of $R_G^{0.4}$ for three data series: (red triangles) in situ observations; (black dots) the reconstructed IMF from the IDV(1d) index, as presented here (with uncertainties in B as shown in Fig. 7); (cyan squares) the modelled IMF since 1845; and (yellow diamonds) the modelled IMF before 1845.

the modelled variation after 1845 all show a linear variation. This appears to show a limiting minimum value or “floor” in the IMF of around 4 nT when the group sunspot number R_G falls to zero. The only data in Fig. 9 which fall below such a floor is the modelled data for before 1845 and so one might be tempted to dismiss these modelled values as an error, even though the modelling is very successful after 1845 (the cyan squares match the reconstructed values well, giving $r = 0.810$, $S > 99.99$). Note that there are unknown uncertainties in the early group sunspot data and these will have added to the scatter in Fig. 9 for the early data.

In this context, it is important to realise that the correlation is only with the simultaneous group sunspot number and that the prior history of the solar activity is not a factor. This is contradicted by a number of pieces of pertinent evidence. Firstly the observed and reconstructed IMF variations clearly show some long-term persistence with the variation over one cycle being an evolution of that seen during the previous cycle. This has been confirmed quantitatively by Lockwood et al. (2011), who show significant autocorrelation (and predictability) over several solar cycles in open solar flux and cosmic ray modulation potential. Another indicator of this persistence is the success of precursor methods, whereby the level of geomagnetic activity, the solar polar field or the near-Earth IMF are used to predict the size of the subsequent maximum in sunspot number (Svalgaard et al., 2005; Lockwood and Owens, 2014). Lastly, we note the importance of the prior history of solar activity in the success of continuity models in explaining the reconstructed OSF variation, as initially demonstrated by Solanki et al. (2000) and as shown by the cyan squares in Fig. 9.

This means that it is not just the contemporaneous sunspot number that determines the open solar flux and near-Earth IMF (and related parameters such as the heliospheric modulation parameter), but its prior history also has an influence. It is not immediately obvious over how long a prior period this history matters, but the success of the precursor methods tells us that the polar solar fields half a solar cycle in advance have an influence and Lockwood et al. (2011) show there is some (non-zero) persistence in sunspot numbers up to two solar cycles in advance and a much stronger persistence in open solar flux and heliospheric modulation potential. This provides an explanation of why the behaviours before and after 1844 are so different in the modelled data in Fig. 9: the data after 1844 containing persistently high sunspot numbers (and a grand solar maximum) whereas before 1844 they contain lower sunspot numbers and both the Dalton and the Maunder minima. Here we demonstrate the effect of prior history by using the mean of the group sunspot number over the previous 11 yr, $\langle R_G \rangle_{11}$, in addition to the simultaneous annual value, R_G . We also make a small second correction to allow for the known effect of the solar wind speed using the theory of the Parker spiral. We use a simple combination of R_G , and $\langle R_G \rangle_{11}$, given by

$$B_p = \alpha R_G^n + \beta \langle R_G \rangle_{11}^m + \gamma. \quad (2)$$

This simple form has been chosen to introduce some dependence on the prior history of the sunspot number (via the term $\beta \langle R_G \rangle_{11}^m$): it is in no way an optimised form and is used here purely for demonstration purposes. Using the Nelder–Mead search method (Nelder and Mead, 1965; Lagarias et al., 1998), the maximum correlation between the predicted B , B_p (given by Eq. 2) and B is obtained for the coefficients $\alpha = 0.462$, $n = 0.454$, $\beta = 0.090$, $m = 0.668$ and $\gamma = 0.006$. Using these coefficients and Eq. (2) to compute B_p , Fig. 10 shows the plot of B as a function of B_p for the same data sets as shown in Fig. 9 (using the same symbols). In addition, Fig. 10 makes allowance for the effect on near-Earth IMF of variations of the solar wind speed, V_{SW} , via Parker's spiral theory. This allowance is relevant because increased solar wind speed causes the spiral to unwind and so lowers the near-Earth IMF; as a result, for a given OSF, a higher V_{SW} gives a lower near-Earth B . The factor $f(V_{SW})$ used in Fig. 10 gives a correction to a constant solar wind speed of V_o and is given by

$$f(V_{SW}) = \sin(\varphi) / \sin(\varphi_o), \quad (3)$$

where the garden-hose angle for solar wind speed, V_{SW} , is $\varphi = \tan^{-1}(V_{SW}/\omega R)$ and for the solar wind speed, V_o , is $\varphi_o = \tan^{-1}(V_o/\omega R_1)$, ω is the angular velocity of the solar corona in the frame of the fixed stars and R_1 is one astronomical unit. The average solar wind velocity over recent cycles (470 km s⁻¹) is used for V_o . This correction for the effect of solar wind speed on the IMF can be made using observed annual means of V_{SW} for 1964 onwards and for

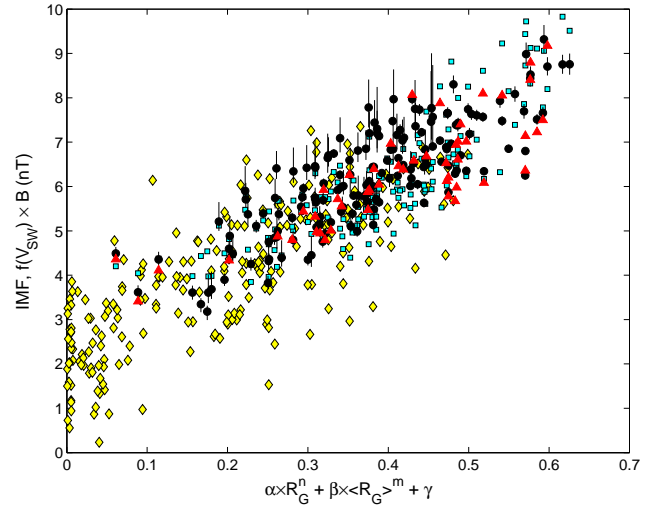


Fig. 10. A modified version of Fig. 9 with $f(V_{SW})B$ shown as a function of $B_p = \alpha \times R_G^n + \beta \times \langle R_G \rangle_{11}^m + \gamma$ (where R_G is the simultaneous annual mean group sunspot number and $\langle R_G \rangle_{11}$ is the mean value over the preceding 11 yr) with best-fit coefficients $\alpha = 0.462$, $n = 0.454$, $\beta = 0.090$, $m = 0.668$ and $\gamma = 0.006$. The factor $f(V_{SW})$ allows for the effect on IMF B of the solar wind speed V_{SW} using Eq. (2), based on Parker's spiral theory.

1868 onwards for the reconstructed data using the solar wind variation derived by Lockwood and Owens (2014), using the combination of the corrected *aa* index, aa_C , and the corrected IDV(1d) index discussed in previous sections of the present paper. For the model, the correction term is unity because the modelled values were generated from OSF predictions and then converted to B assuming constant solar wind speed.

Figure 10 shows that even using this simple form of Eq. (2) the same linear relation between B and B_p is shared by all three data sets. The correlation coefficient for the 167 annual means of the geomagnetic reconstruction of B in Fig. 9 is $r = 0.835$, whereas for that the 144 annual means in Fig. 10 is $r = 0.833$. Allowing for the number of fit variables, the Fisher Z test gives that the difference between these two correlations is significant at just 4.5 %. For the 49 observed IMF annual means the correlations are 0.837 and 0.843, respectively, a difference that is significant at only the 5.5 % level. However, for the 399 modelled data points the correlation rose from 0.729 to 0.861 which is a significant increase at over the 99.99 % level. Furthermore, all data sets are now giving very similar correlations. Note that the dependence of the term in $\langle R_G \rangle_{11}$ is relatively weak with the coefficient β being roughly a fifth of the corresponding coefficient α for R_G and the exponents $n = 0.454$ and $m = 0.668$, meaning that the term αR_G^n is typically twice $\beta \langle R_G \rangle_{11}^m$.

Thus, making a relatively small and simple allowance for the prior history of the sunspot number explains how the modelled values can be below the apparent floor value before 1845, whilst matching the data and reconstruction

after 1845. Making this simple allowance for the prior R_G history makes no significant difference to the fits to data after 1845 (for which we have geomagnetic reconstructions) or after 1964 (for which we have in situ observations). This directly demonstrates that the available IMF data and reconstruction do not extend into a wide enough range of conditions to differentiate between the two sunspot number fits and hence allow detection of the dependence on the prior history of R_G . Thus one has to look very carefully at the assumptions inherent in a linear regression fit before using it to derive an estimate of any floor value to the near-Earth IMF.

8 Conclusions

We find that the correction of Svalgaard (2014) to IDV(1d) for solar cycle 11 is largely correct. We do this in three ways, firstly using daily $Ak(H)$ indices and comparing to the corrected aa index, aa_C . Secondly we use the newly digitised St Petersburg data which is conveniently close to Helsinki and so provides a very valuable test. Finally we derive IDV(1d)_{HLS,D} from the declination observations at Helsinki. We show that the $Ak(H)$ test and recalibration is roughly twice as accurate as methods using the average diurnal range, comparing to other stations or to group sunspot number. The strong dependence of diurnal range-based proxies on solar wind speed means that, as yet, we do not have any digitised geomagnetic data on which we can reconstruct the IMF before the start of the Helsinki data midway through 1844. Thus, accurate reconstruction between 1835 (just 3 yr after Gauss constructed the first magnetometer) and 1844 is not possible until we have found, digitised and calibrated, some other historic magnetometer data. In relation to that task, we urge considerable caution in using average diurnal range and sunspot number to calibrate and correct those data. Our tests show that such recalibrations are only accurate to $\pm 20\%$ even on modern magnetometer and group sunspot data and led to Bartels' u index (and hence the IDV index that employed it) becoming incorrectly similar to the sunspot number in the years before 1872.

We have shown that there is a consistent behaviour of observed, reconstructed and modelled IMF variations only if the role of prior solar activity over an extended period is considered, and not just the simultaneous value. This analysis indicates that estimates giving a high-level floor to the near-Earth IMF arise from the use of data that does not extend back into the Dalton or Maunder minima and fail to recognise the role of prior solar activity. One important implication of the role of the prior history is that it means predictions of future solar activity have a sound physical basis in addition to the statistical predictability found in heliospheric parameters (such as open solar flux, the interplanetary magnetic field and the cosmic ray modulation potential) over several solar cycles (Lockwood et al., 2010). These predictions have ranged from half a cycle in advance (Svalgaard et al., 2005)

to several cycles in advance (Lockwood, 2010; Barnard et al., 2011). The applications of such predictions in the area of space weather are well documented and include satellite orbit prediction, integrated radiation dose prediction for electronics on spacecraft and aircraft, corrosion maintenance planning on long pipelines, and many more. Although the influence of future changes in solar activity on global mean temperatures is predicted to be small (Feulner and Rahmstorf, 2010; Jones et al., 2012) there are interesting indications of some regional influences, particularly in winter (Lockwood et al., 2010; Lockwood, 2012; Woollings et al., 2010), increasing the importance of making predictions of the level of solar activity.

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Appendix A

Glossary of indices and acronyms used

aa index. Three-hourly indices developed by Mayaud (1971) and derived from the k values at two near-antipodal stations: one in southern England, one in Australia. In both hemispheres, three different stations were needed to give a continuous index: in the north they are Greenwich (1868–1925), Abinger (1926–1956), and Hartland (1957–present) and in the south they are Melbourne (1868–1919), Toolangi (1920–1979), and Canberra (1980–present). This yields the aa_N and aa_S indices for the Northern and Southern hemispheres and aa is defined as the arithmetic mean of the two.

aa_C index. A corrected version of the aa index that allows for a putative calibration error in 1957 associated with a move of the Northern Hemisphere's aa station from Abinger to Hartland. The correction to aa to yield aa_C is discussed further in Paper 4 (Lockwood et al., 2014).

$ak(H)$ index. A three-hourly “equivalent amplitude” index of geomagnetic activity for a specific station or network of stations expressing the range of disturbance in the horizontal magnetic field, H : $ak(H)$ is scaled from the standard three-hourly range k indices such that it equals 0, 3, 7, 15, 27, 48, 80, 140, 240, and 400 for $k = 0$ –9, respectively. Multiplying by constant factor c for each station converts into nT where c equals the lower limit of the range for the $k = 9$ band, divided by 250.

$Ak(H)$ index. A daily index of geomagnetic activity for a specific station or network of stations, being the average of the eight, three-hourly $ak(H)$ indices for that day. The $Ak(H)$ for Helsinki, $Ak(H)_{HEL}$, was used to extend the aa index series from 1868 back to 1844 by Nevanlinna and Ketola (1993).

$Ak(D)$ index. The same as $Ak(H)$, but instead of being based on the variation range of the horizontal component H , it is based on that in the field declination, D (the angle between geographic and geomagnetic north).

Diurnal range, ΔH . The difference between the maximum and the minimum of the horizontal field component H detected at one station in 1 day.

Group sunspot number, R_G . Originally generated by Hoyt and Schatten (1998) and extending back to 1610, this sunspot index is based on the number of sunspot groups G on the visible disc of the Sun, scaled to match the international sunspot number in later years. Here we deploy the corrections proposed by Usoskin et al. (2003) and Vaquero et al. (2011).

IDV index. The interdiurnal variation index introduced by Svalgaard and Cliver (2005) based on Bartels' u index. The main difference between IDV and u is that instead of using daily mean values of H , the hourly mean (or spot value) closest to solar local midnight is employed as a means of eliminating the diurnal variation. As for u , the difference between values on successive days is taken. The geomagnetic latitude

(Λ) normalisation is also slightly different, using an empirically derived $1/\cos^{0.7}(\Lambda)$ dependence.

IDV(1d) index. A variant of the IDV index introduced in Paper 1, which returns to using daily means, as used by Bartels to generate the u index, which gives suppression of noise. It is referred to as IDV(1d) rather than u as it uses stations in a band of geomagnetic latitude Λ chosen to ensure the index has no dependence on solar wind V_{SW} (in particular it avoids low latitude stations for which there is a dependence on $V_{SW}^{-0.4}$).

IDV(1d)_D test values. These are derived in the same way as the standard IDV(1d) index but using the declination data rather than H . The results, for example, for Helsinki station are termed IDV(1d)_{HLS,D}.

IMF: interplanetary magnetic field. Unless otherwise stated, this refers to the near-Earth field.

International sunspot number, R . (Also known as the Wolf number, the relative sunspot number, or the Zürich number). This is defined as $c(10G + N)$, where G is the number of sunspot groups on the visible disc of the Sun, N is number of individual spots, and c is an observer calibration factor that varies with location and instrumentation. Here we use the data series extending back to 1749 published by the Solar Influences Data Analysis Center (SIDC), Belgium.

k index. Three-hourly indices introduced by Julius Bartels in 1938 to quantify disturbances in the horizontal component of Earth's magnetic field (H) at a given station with an integer between 0 (quiet) and 9 (severely disturbed). The k values are determined by the range of variation in 3 h intervals, after subtraction of the quiet-day variation, using a quasi-logarithmic set of thresholds specific to each station to normalise them to the values seen at the Niemegk station.

u index. The u index was developed by Bartels (1932). It was based on the absolute value of the difference between the mean values of H for a day and for the preceding day (an effective way of removing quiet-time variation). The u index is the weighted mean of data from a collection of stations. Prior to averaging the data from the various stations, each was normalised to the magnetic latitude (Λ) of Niemegk using an empirical $1/\cos(\Lambda)$ dependence.